



### Direct Variational Assimilation of Radar Reflectivity and Radial Velocity Data: Issues with Nonlinear Reflectivity Operator and Solutions

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Liu, Xue and Kong (2018, to be submitted)

# Outline

- Background and motivations
- Reflectivity observation operator
- Problems and treatments associated with the reflectivity operator
- Experimental results
- Summary

## Previous studies of Z assimilation

➢Cloud analysis (Albers et al. 1996; Xue et al. 2003; Hu et al. 2006; Kain et al. 2010; Sun et al. 2014)

- Relatively cheaper computational cost and easier implementation
- Rely on empirical relationship
- Does not use error statistics

➢EnKF (Aksoy et al. 2009, 2010; Jung et al. 2008a; Jung et al. 2008b; Tong and Xue 2005; Xue et al. 2006)

- Does not use TL and Adj.
- Can include nonlinear operator but the solution is suboptimal (Lorenc 2003)
- Affected by Sampling error

## Previous studies of Z assimilation

#### ➤Variational

- 4DVar (Sun and Crook 1997)
  - Compared direct assimilation of Z and assimilation of derived rainwater
  - Considered warm rain only
- 3DVar (Xiao et. al. 2007; Gao and Stensrud 2012)
  - Has low computational cost
  - Uses flow-independent **B**
- Hybrid EnVar (Carley 2012; Wang and Wang 2017; Kong et al. 2018)
  - Using logarithmic mixing ratios as the control variables (Carley 2012)
  - a pure EnVar method without using the TL and ADJ Z operator (Wang and Wang 2017)
  - Compared direct Z assimilation using hybrid EnVar and EnKF (Kong et al. 2018)

### Motivations

- Aim to further improve direct Z DA based on the variational framework
  - Study the problems with the nonlinear Z operator when considering Multi-phased MP
  - Propose some treatments to deal with the problems
  - Examine the impacts of the treatments through OSSEs.

#### Z operator considering multi-phased hydrometeors (take Lin scheme as example)

$$Z = 10 \log Z_e \qquad \qquad Z_e = Z_{er} + Z_{es} + Z_{eh}$$

$$Z_{er} = 3.63 \times 10^9 \times (\rho q_r)^{1.75} \quad Z_{es} = \begin{cases} 9.80 \times 10^8 \times (\rho q_s)^{1.75} & T_b \le 0^\circ \\ 4.26 \times 10^{11} \times (\rho q_s)^{1.75} & T_b > 0^\circ \end{cases} \qquad Z_{eh} = 4.33 \times 10^{10} \times (\rho q_h)^{1.75}$$

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- Background and motivations
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- Problems and treatments with the reflectivity operator
  - Using mixing ratio as control variable (CVq)
  - Using logarithmic mixing ratio as control variable (CVlogq)
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#### Problems with CVq

$$\nabla J = \mathbf{B}^{-1}(\mathbf{x} - \mathbf{x}_b) + \mathbf{H}^{\mathrm{T}} \mathbf{R}^{-1}(H(\mathbf{x}) - \mathbf{y})$$

Single-phase TL:  $H = \frac{\partial Z}{\partial q_r} = \frac{17.5}{\ln 10 \times q_r}$  Multi-phase TL:  $H = \frac{\partial Z}{\partial q_r} = \frac{6.35 \times 10^{10} \times \rho^{1.75} q_r^{-0.75}}{\ln 10 \times Z_e}$ 

- For single-phased hydrometeor,  $q_r \Rightarrow 0, \partial Z / \partial q_r \Rightarrow \infty$ a lower limit for q is used (Sun and Crook 1997, Wang et al 2013, Wang and Wang 2017).
- For multi-phased hydrometeors,  $Z_e \Rightarrow 0, \partial Z / \partial q_r \Rightarrow \infty$ a new treatment of modified  $Z_e$  (instead of q) is proposed

## Special Treatments for CVq

Old treatment:

adding a lower limit to the mixing ratios (qLim, Sun and Crook 1997)

$$Z = H(\max(q, \varepsilon_q)) \qquad \frac{\partial Z}{\partial q} = \begin{cases} \frac{\partial Z(q)}{\partial q} & q > \varepsilon_q \\ 0 & q \le \varepsilon_q \end{cases}$$

\*Limitation: Reflectivity observations are ignored when  $q \leq \varepsilon_q$ , due to zero TL term.

#### ► New treatments:

1. Modify  $Z_e$  if smaller than  $\mathcal{E}_{Z_e}$  (ZeLim)

$$Z_e = \begin{cases} Z_e(q) & Z_e \ge \varepsilon_{Z_e} \\ Z_e(q) + \varepsilon_{Z_e} & Z_e < \varepsilon_{Z_e} \end{cases}$$

2. Use a separate pass to assimilate Vr data (VrPass)

#### Problems with logarithmic transformation in CVlogq

Due to the nonlinear logarithmic transformation between q and log(q), the effective spatial correlations of background error in q space is not Gaussian-like.

- $P_{ij}$ : log(q) background error correlation in q space
- $\hat{\rho}_{ij}$  : log(q) background error correlation in log(q) space (Gaussian like)

- 1. When background is homogeneous , and relatively larger than analysis increment
- 2. When background is homogeneous , and relatively smaller than analysis increment
- 3. When background is inhomogeneous , and relatively smaller than analysis increment

$$\boldsymbol{\rho}_{ij} = \frac{\left[1 + \sum_{k=2}^{\infty} (\ln 10)^{k-1} \delta \hat{q}_i^{k-1} (k!)^{-1}\right]}{\left[1 + \sum_{k=2}^{\infty} (\ln 10)^{k-1} \delta \hat{q}_j^{k-1} (k!)^{-1}\right]} \cdot \frac{q_{b_i}}{q_{b_j}} \frac{\delta q_i}{\delta q_j} \hat{\rho}_{ij}$$

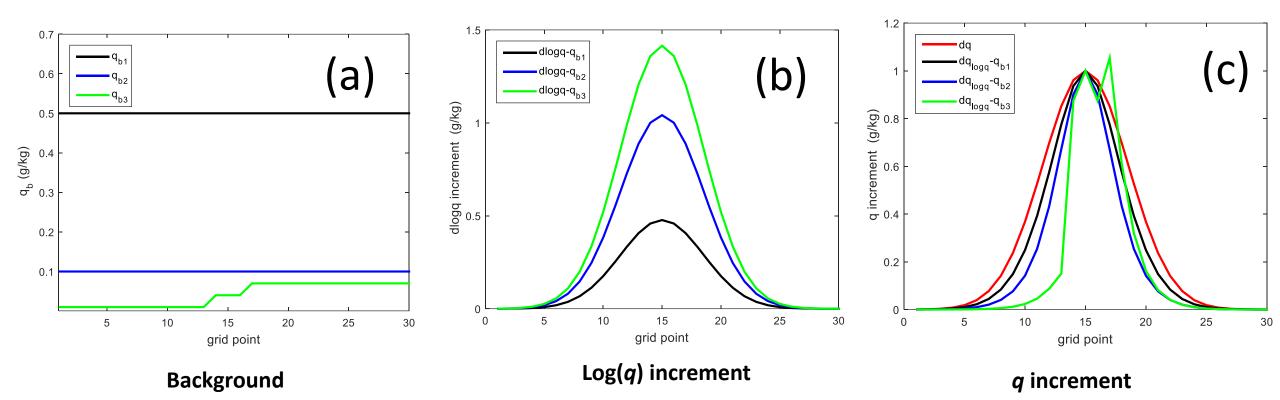
$$\rho_{ij} \approx \hat{\rho}_{ij}$$

$$\rho_{ij} \leq \hat{\rho}_{ij}$$

$$\rho_{ij} \approx \frac{q_{b_i}}{q_{b_j}} \cdot \hat{\rho}_{ij} = \begin{cases} > \hat{\rho}_{ij} & q_{b_i} > q_{b_j} \\ < \hat{\rho}_{ij} & q_{b_i} < q_{b_j} \end{cases}$$

### Single observation test

- > Assumptions:
  - Gaussian correlation functions with the same decorrelation scales are used for CVq and CVlogq
  - Same innovations (O-B)
  - Three different background values



- When background is homogeneous, and relatively larger than analysis increment (black line in Fig. a), CVlogq analysis (black line in Fig. c) is closed to CVq analysis (red line in Fig. c)
- When background is homogeneous, and relatively smaller than analysis increment (blue line in Fig. a), the spread of analysis increment using CVlogq becomes narrow (blue line in Fig. c)
- When background is inhomogeneous, and relatively smaller than analysis increment (green line in Fig. a), the spread of analysis increment using CVlogq becomes non-Guassian (green line in Fig. c).

# Special Treatment for CVlogq

• A lower limit is added to background when converting  $\delta q$  back to  $\delta q$  (X<sub>b</sub>Lim)

 $\hat{q} = \log(q)$ 

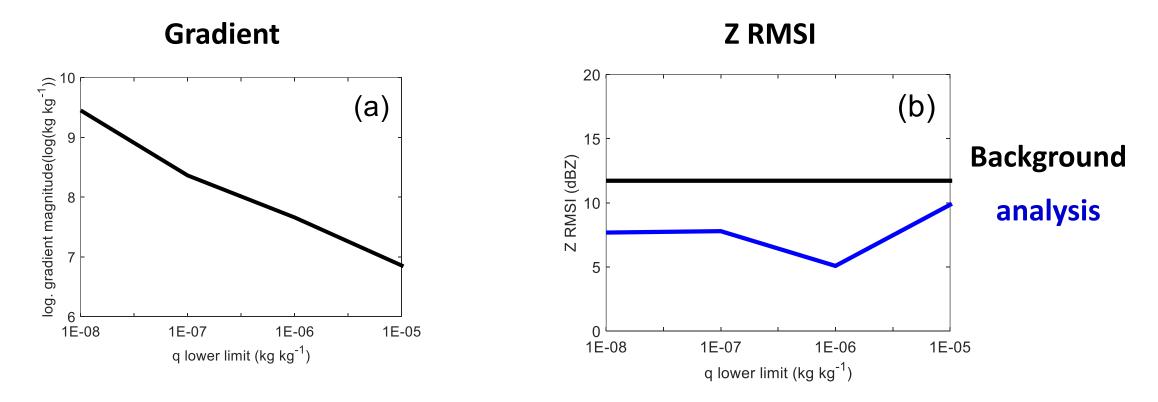
$$\delta q = 10^{\delta q + \log(\max(q_b, \varepsilon_b))} - \max(q_b, \varepsilon_b)$$

#### Examination of the treatments through OSSEs

- Tested with simulated data from a supercell storm of 20 May 1977 near Del City, Oklahoma
- ARPS-3DVar system and ARPS model
- 2 km horizontal and 0.5 km vertical grid spacings

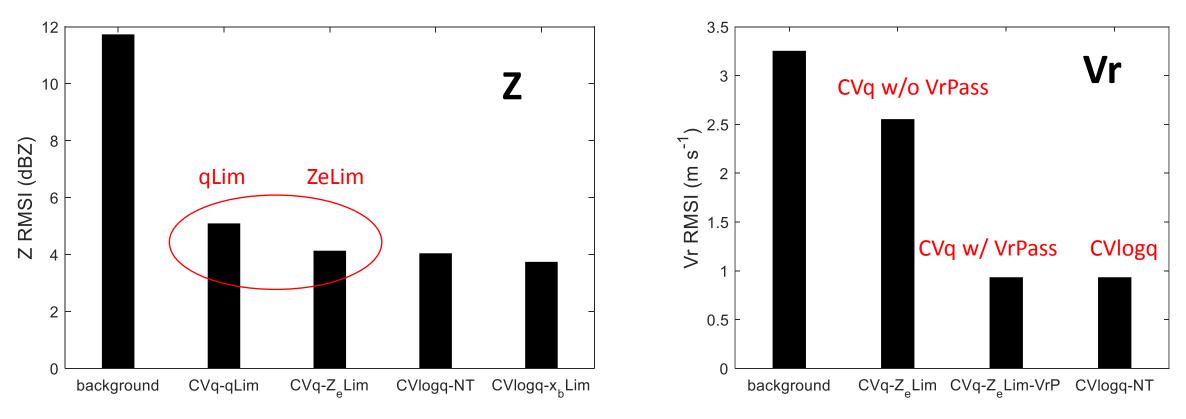
			CVlogq			
treatment	qLim	Z <sub>e</sub> Lim	qLim+VrPass	ZeLim+VrPass	No Treatment	X <sub>b</sub> Lim
Exp. name	CVq-qLim	CVq-Z <sub>e</sub> Lim	CVq-qLim-VrP	CVq-ZeLim-VrP	CVlogq-NT	CVlogq-X <sub>b</sub> Lim

## Optimal threshold for qLim treatment



- The magnitude of the gradient significantly decreases when increasing the qLim threshold.
- 10<sup>-6</sup> kg kg<sup>-1</sup> is the optimal threshold of hydrometeor mixing ratio in this study.

## RMSI for Z and Vr



- Z<sub>e</sub>Lim is better than qLim in terms of a smaller RMSI
- Vr data is difficult to assimilate when Z and Vr data are assimilated together using CVq.
- the Vr RMSIs of CVq with VrPass treatment is similar to CVlogq

## Convergence speeds of using CVq and CVlogq

30

0

20

40

60

80

100

iteration steps

120

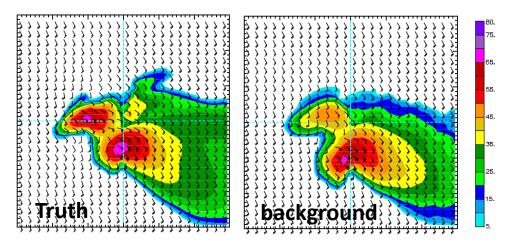
140

160

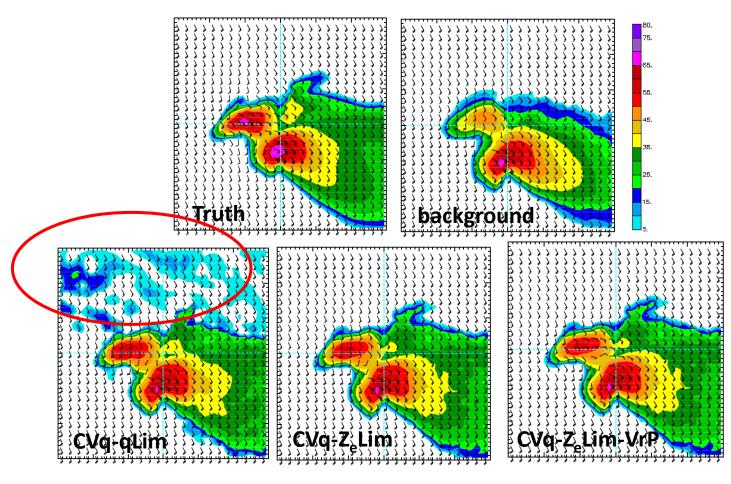
180

200

CVq-qLim CVq-ZeLim 25 CVlogg-NT ICVlogq-x<sub>b</sub>Lim CVlogq converges much faster than CVq. <sup>20</sup> <sup>1</sup> <sup>1</sup> Reason: **N** 10 the gradient range : CVq:  $-10^8$  to  $10^8$  CVlogq:  $-10^2$  to  $10^2$ 5

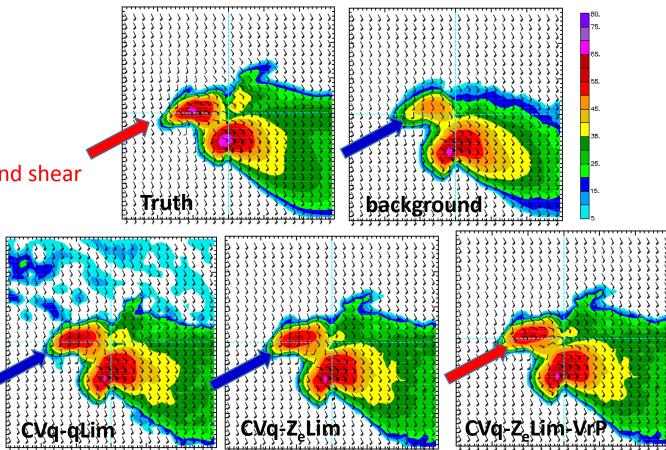


• Applying qLim can introduce spurious Z analyses.

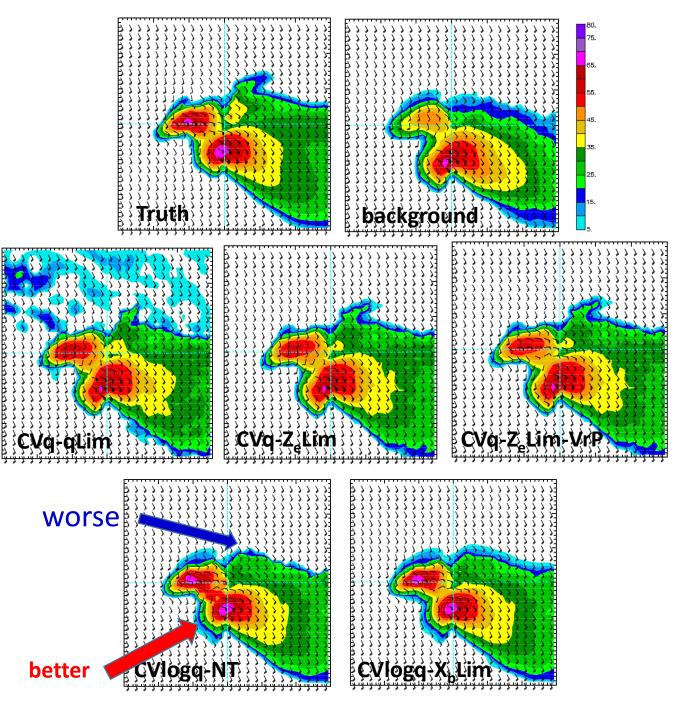


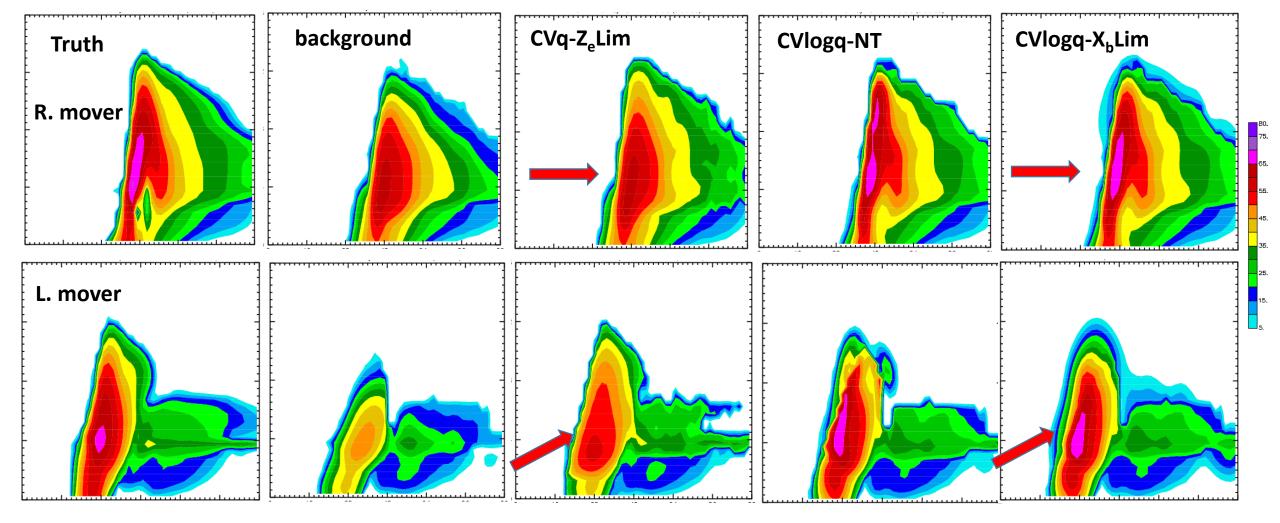
- Applying qLim can introduce spurious Z analyses. Low-level Wind shear
- Using different pass for radial velocity DA, the analysis is greatly improved.

Low-level wind shear

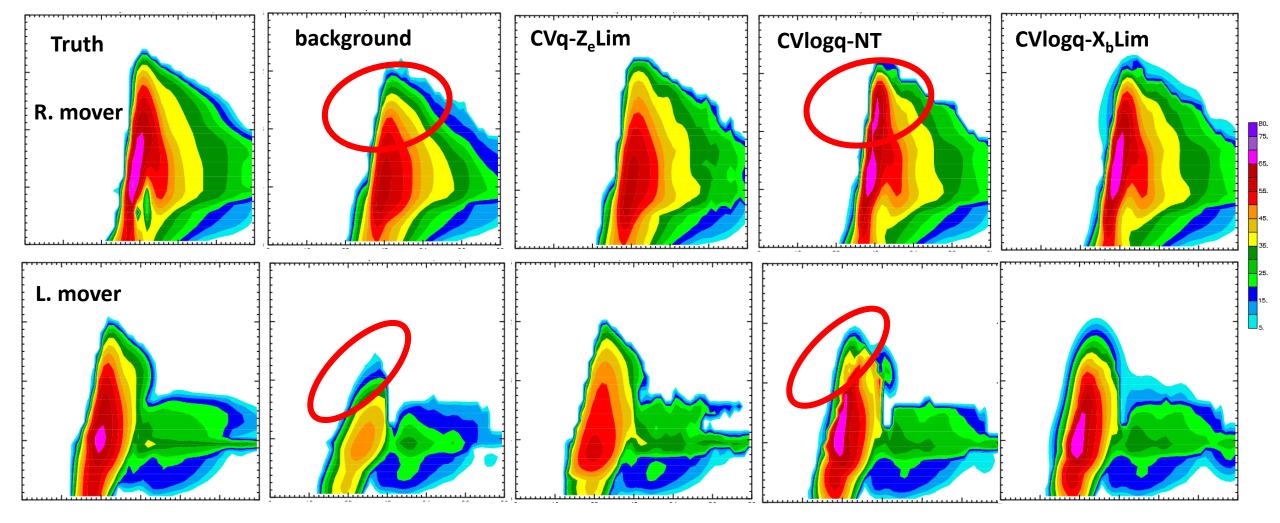


- Applying qLim can introduce spurious Z analyses.
- Using different pass for radial velocity DA, the analysis is greatly improved.
- CVlogq outperforms CVq for analysis of the reflectivity core, but underperforms for analysis of small reflectivity



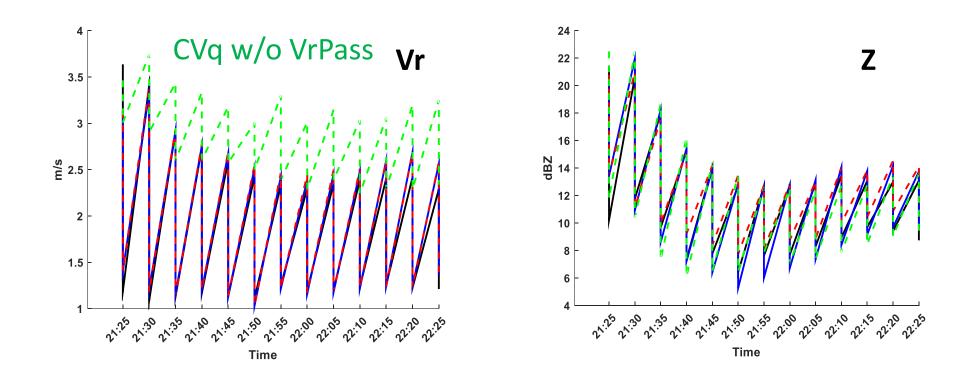


• CVlogq performs betters for the analysis of the high reflectivity cores



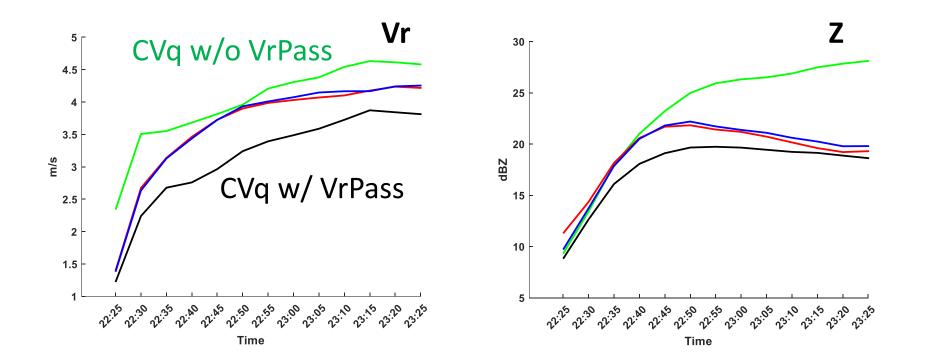
- CVlogq performs betters for the analysis of the high reflectivity cores.
- X<sub>b</sub>Lim treatment can reduce the improper spreading of the analysis increment for CVlogq.

#### RMSI of one hour cycle

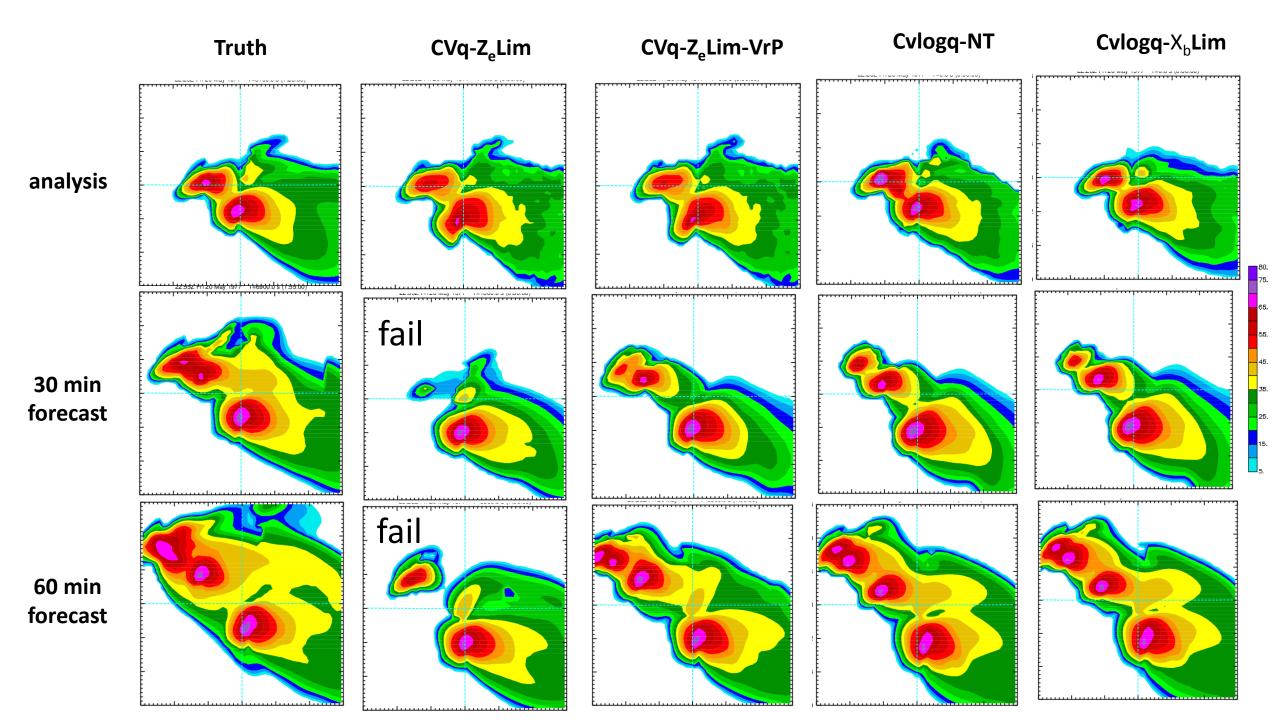


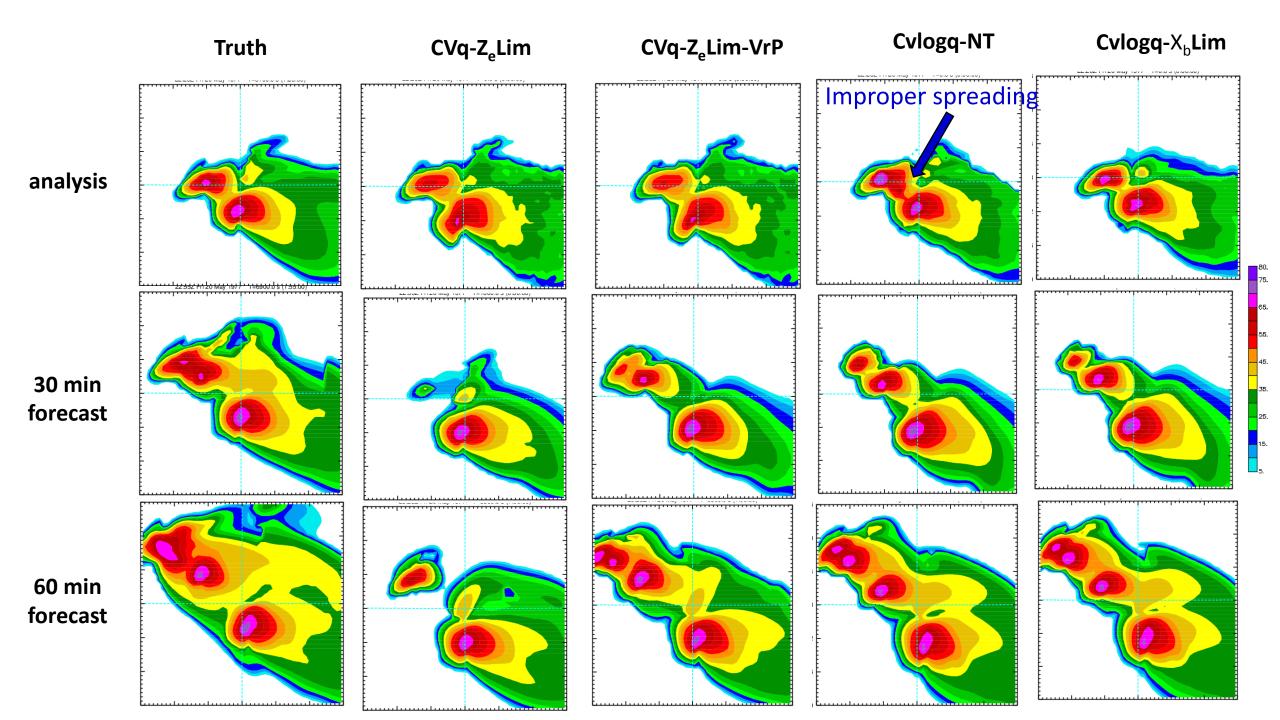
• Green: CVq-Z<sub>e</sub>Lim Black: CVq-Z<sub>e</sub>Lim-VrP Red: Cvlogq-NT Blue: Cvlogq-X<sub>b</sub>Lim

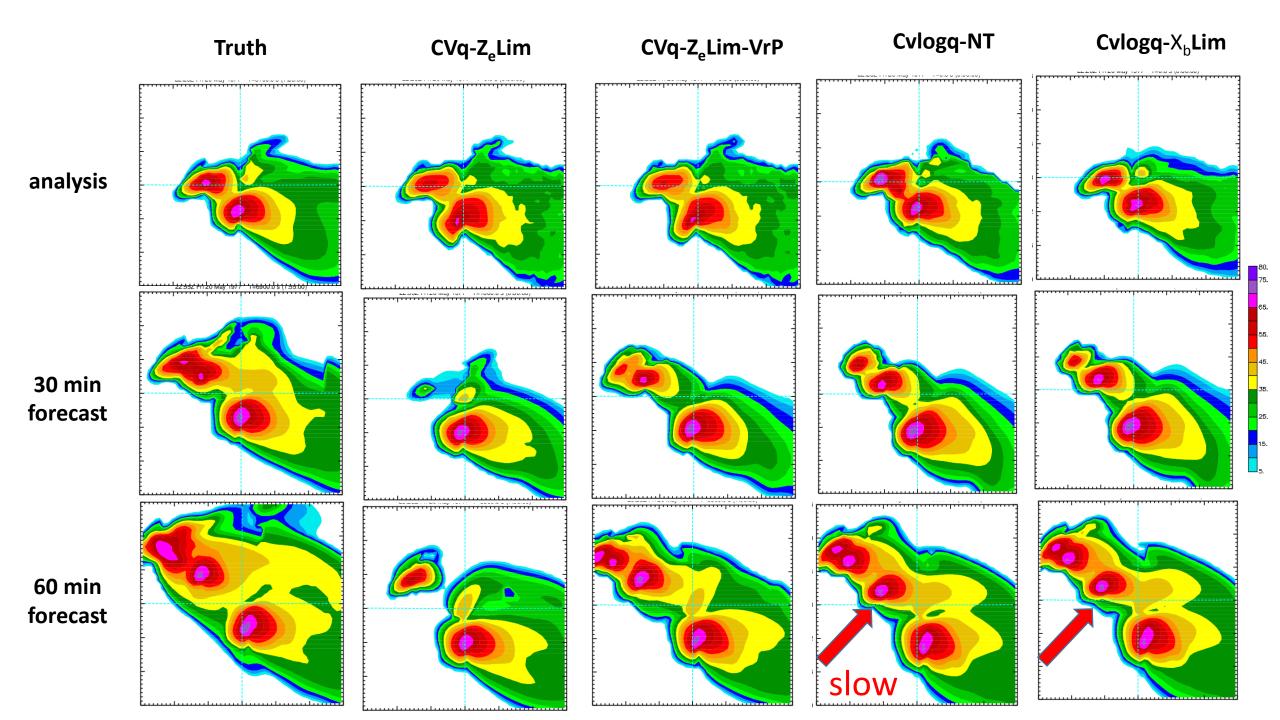
#### RMSEs for one hour forecast



Green: CVq-Z<sub>e</sub>Lim Black: CVq-Z<sub>e</sub>Lim-VrP Red: Cvlogq-NT Blue: Cvlogq-X<sub>b</sub>Lim







# Summary

	Extremely Large gradient	Lacking adjustment for Vr analysis for Vr+Z DA	Improper spreading of Z analysis increment	Lacking adjustment for Z analysis
CVq	Yes, need qLim/Z <sub>e</sub> Lim, Z <sub>e</sub> Lim better	Yes, need VrPass	No	Yes, for large background value (i.e. reflectivity core)
CVlogq	No	No	Yes. Need X <sub>b</sub> Lim	Yes, for small background value

## Ongoing research

#### • Apply the proposed treatments to an EnVar DA system.

