

Waves to Weather

Project A1: “Upscale impact of diabatic processes from convective to near-hemispheric scale”

An estimation of intrinsic limits of predictability using ICON and a stochastic convection scheme

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Outline

Questions

- What is the **intrinsic limit of predictability** that is imposed by the convection?
- What is the **relevance** of this limit for nowadays weather prediction systems?
- How much room is there for further **improvement**?

Outline

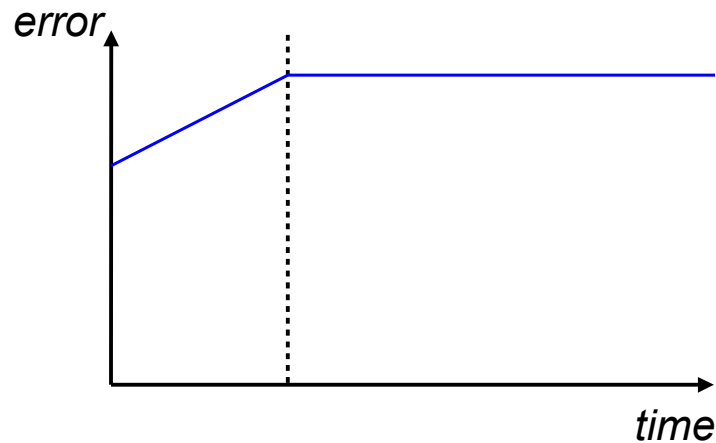
- Introduction
- Experimental setup
- Results

Introduction

Practical and intrinsic predictability

Practical predictability

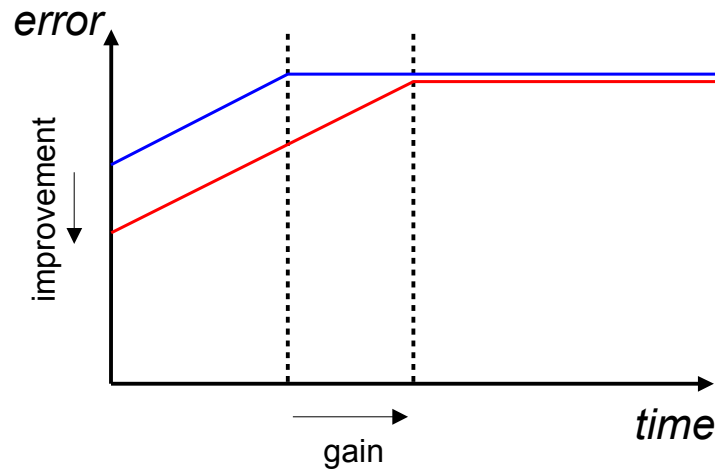
Limit of prediction with currently available models and procedures



Practical and intrinsic predictability

Practical predictability

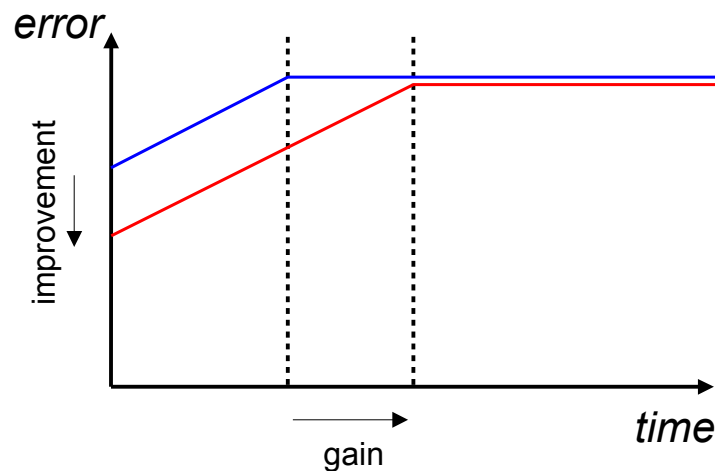
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Practical and intrinsic predictability

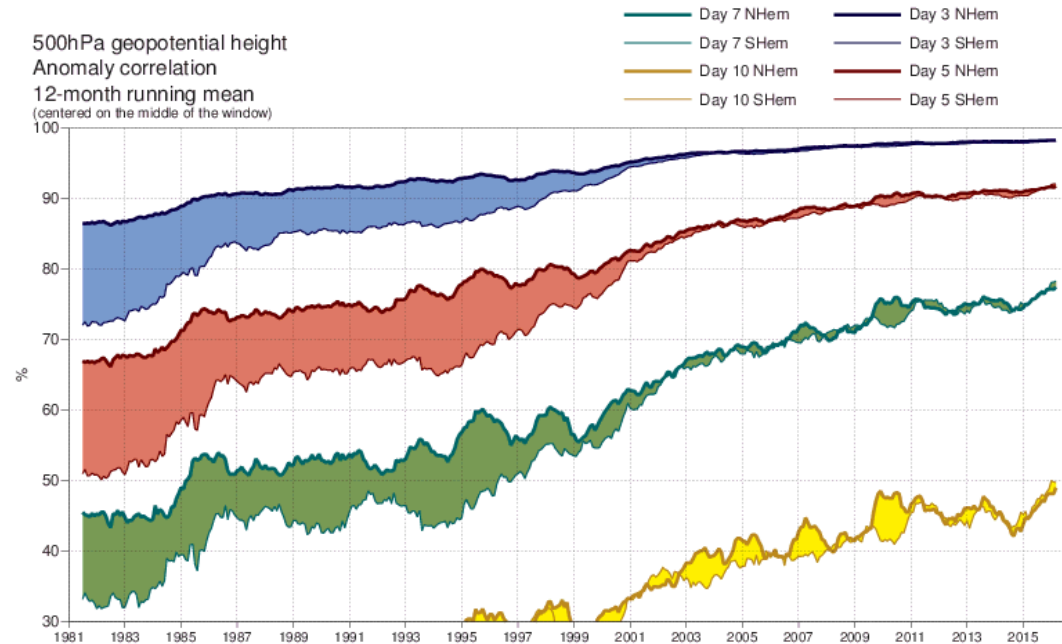
Practical predictability

Limit of prediction with currently available models and procedures



ECMWF forecast skill

500hPa geopotential height
Anomaly correlation
12-month running mean
(centered on the middle of the window)

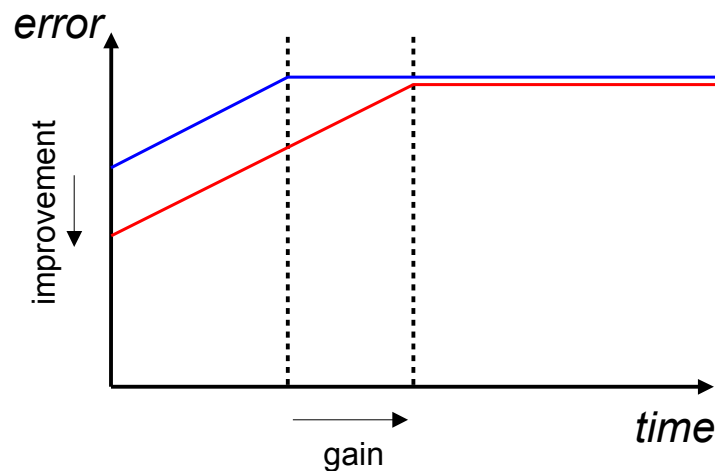


Improvement: 1 forecast-day per decade

Practical and intrinsic predictability

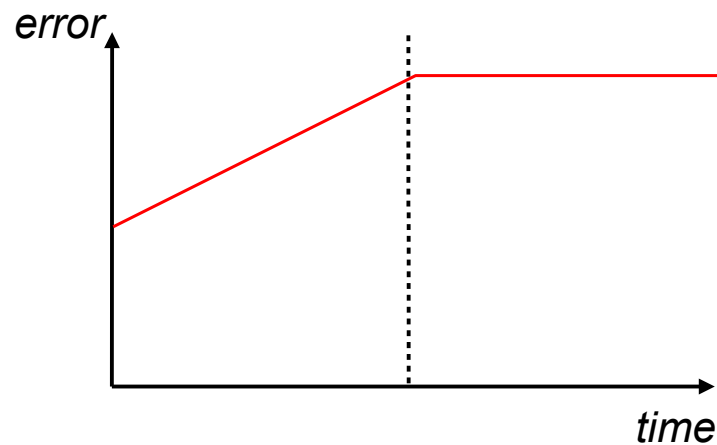
Practical predictability

Limit of prediction with currently available models and procedures



Intrinsic predictability

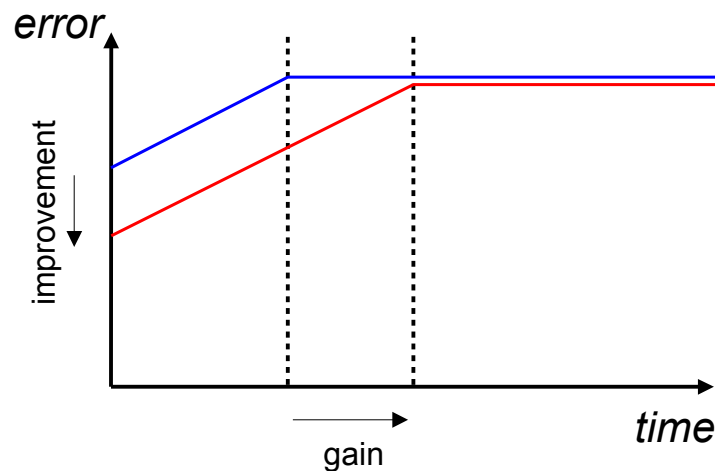
Limit of prediction with perfect procedures and knowledge of the initial state



Practical and intrinsic predictability

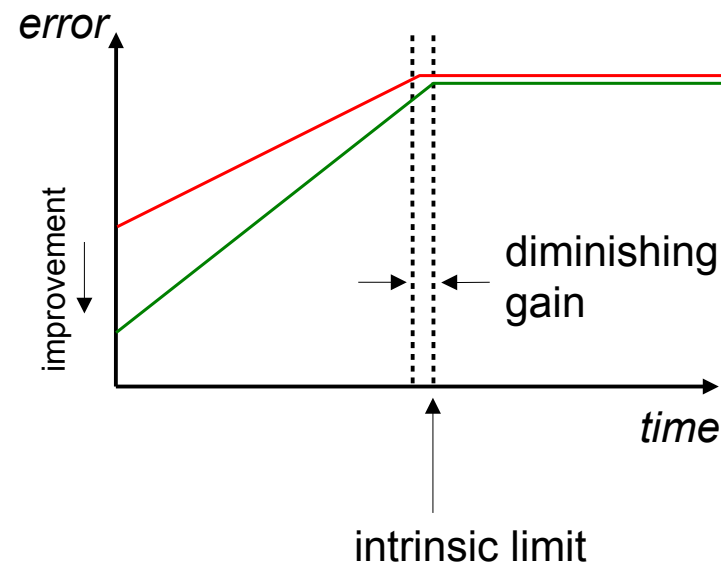
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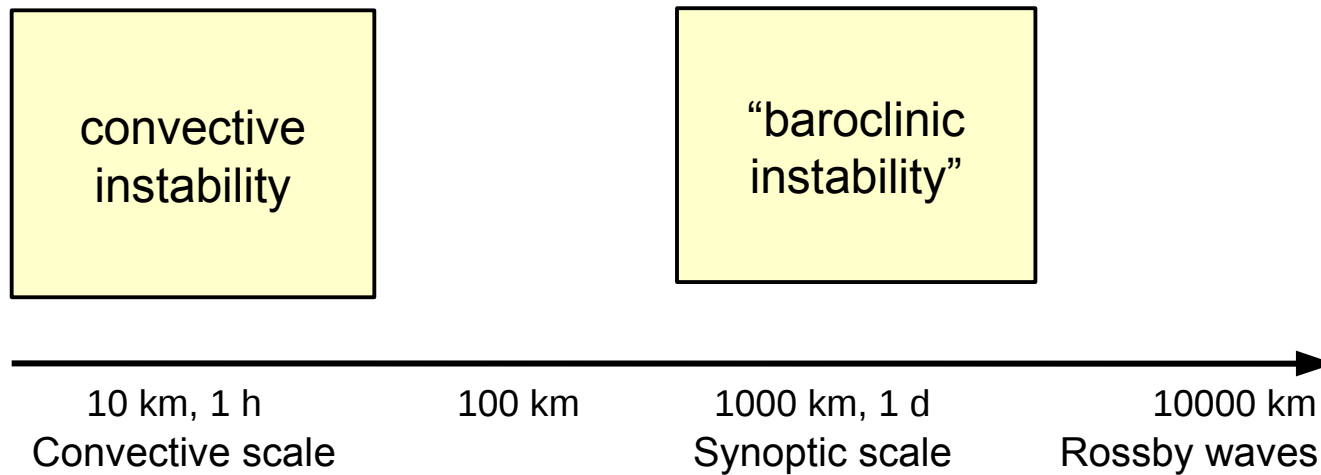


Intrinsic predictability

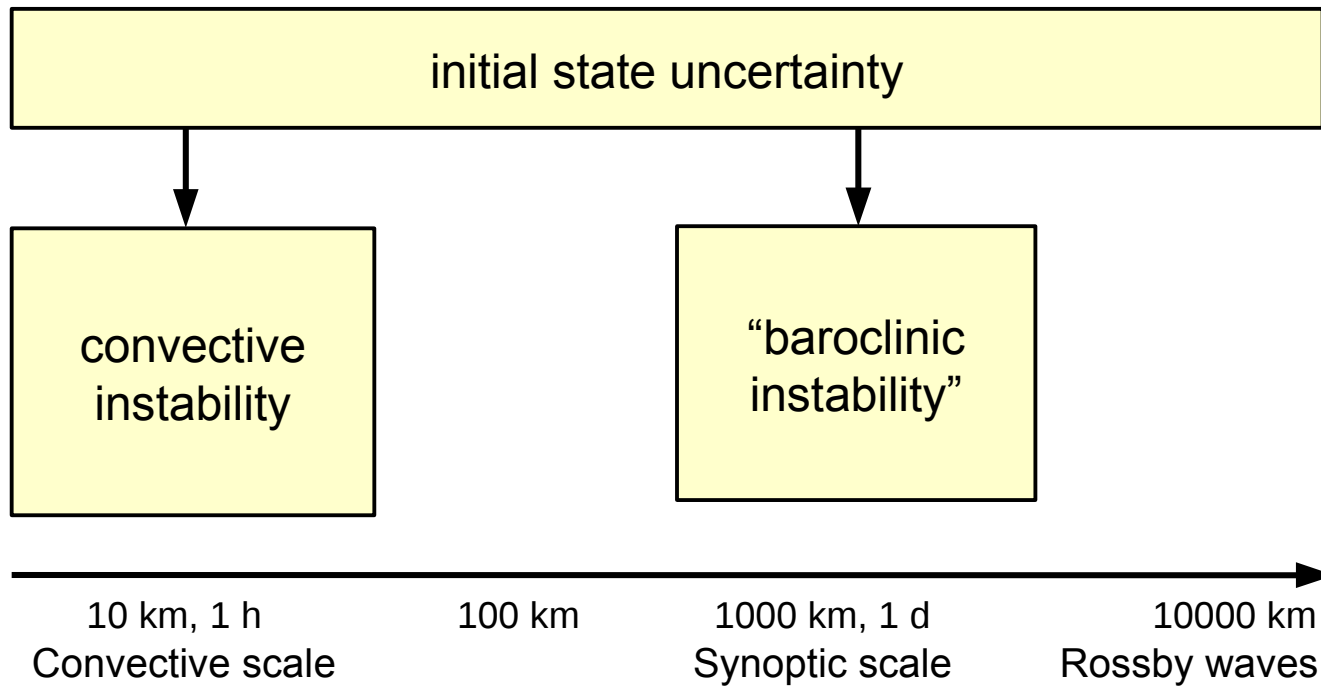
Limit of prediction with perfect procedures and knowledge of the initial state



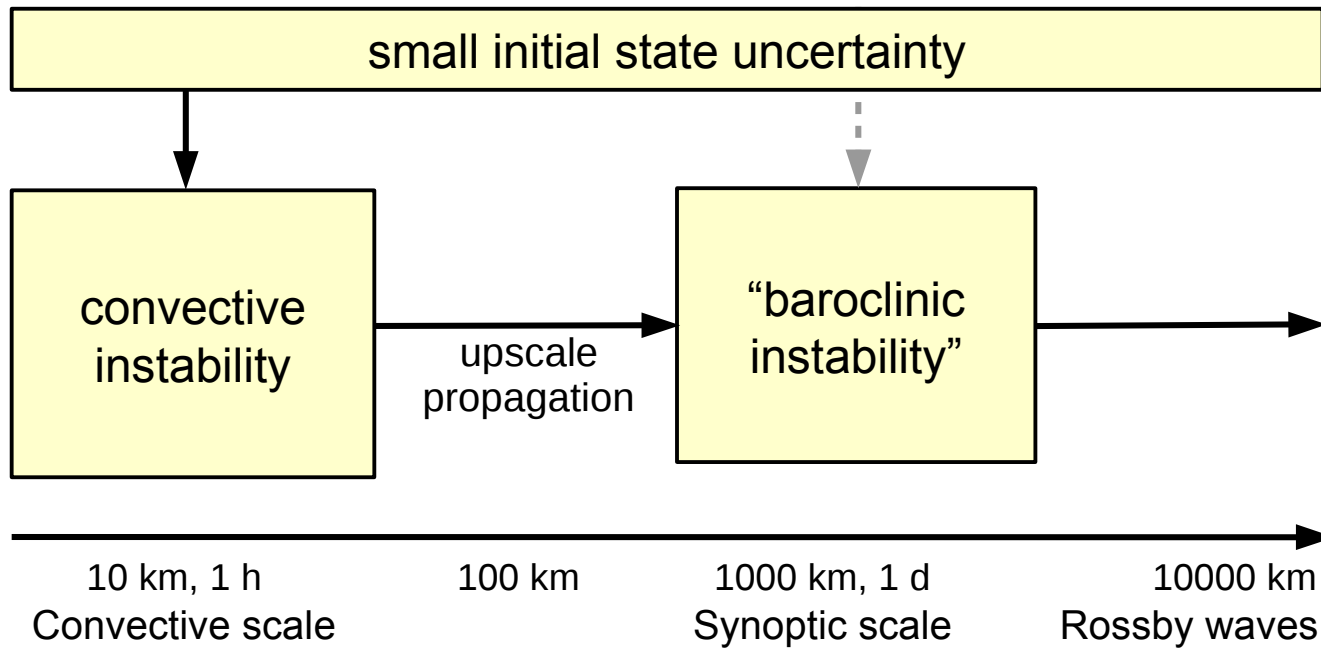
Intrinsic limit of predictability via scale interaction



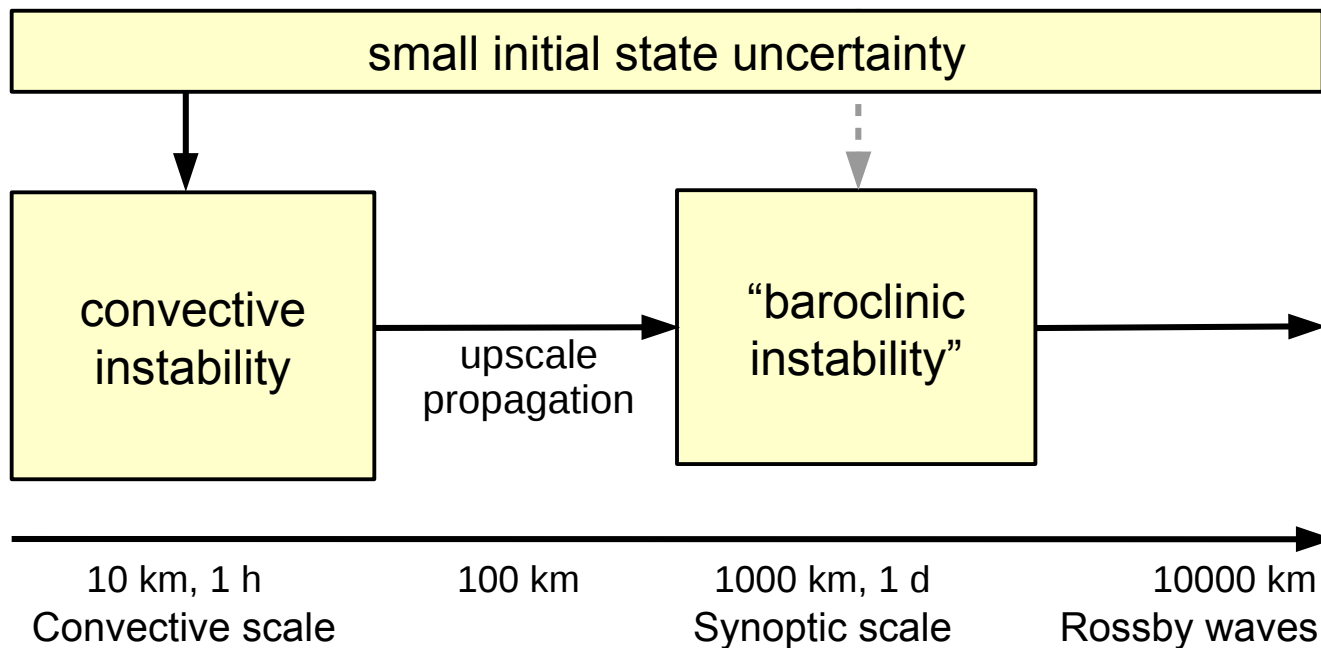
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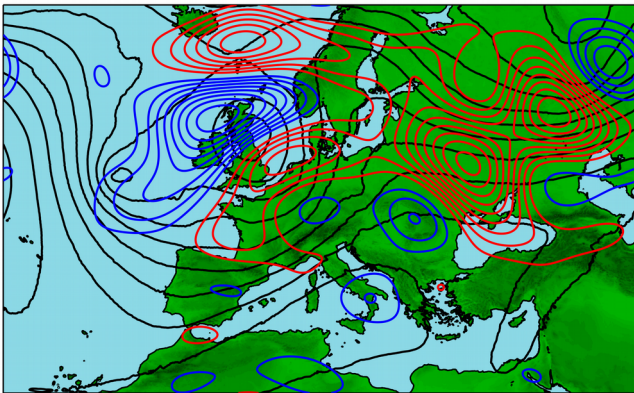
Intrinsic limit of predictability via scale interaction



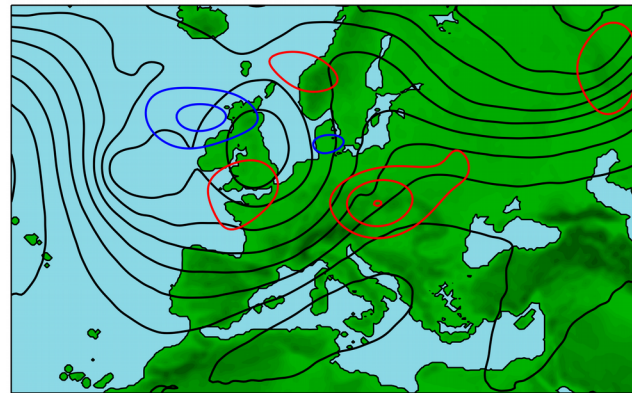
- Quick amplification ($\approx 1\text{h}$) of errors at **convective scale** and subsequent upscale propagation sets the **intrinsic limit** of predictability (Lorenz 1969, Sun and Zhang, 2016)
- Estimate requires a **global model** with an accurate representation of **convection**, but CRM is too expensive
- Is a coarser resolution and a **convection scheme** good enough?

Error growth case study with COSMO (Selz and Craig, 2015a+b)

2.8 km resolution,
no conv-scheme



28 km resolution,
Tiedtke conv.

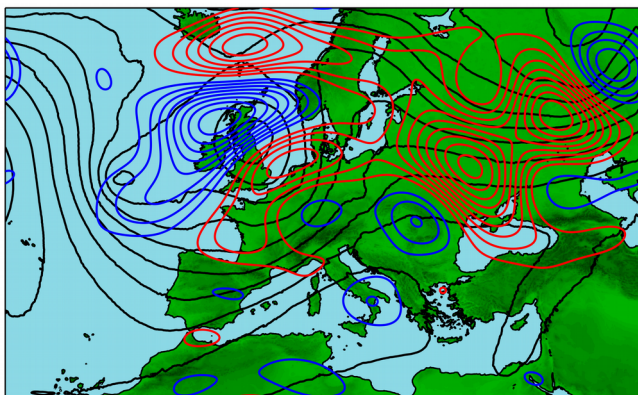


- Conventional convection schemes do not amplify errors near the convective scale sufficiently → **overconfidence**

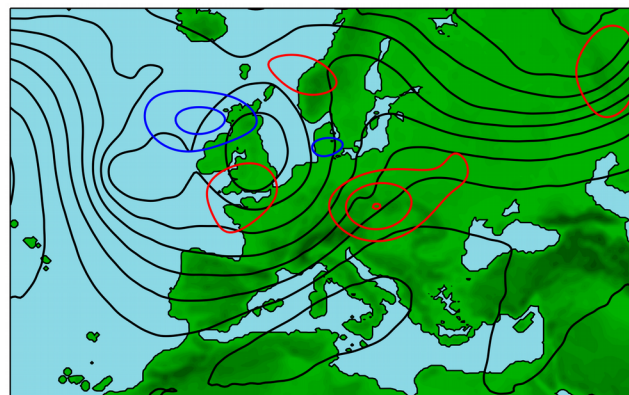
Errors in 500hPa geopot after 60h (color)

Error growth case study with COSMO (Selz and Craig, 2015a+b)

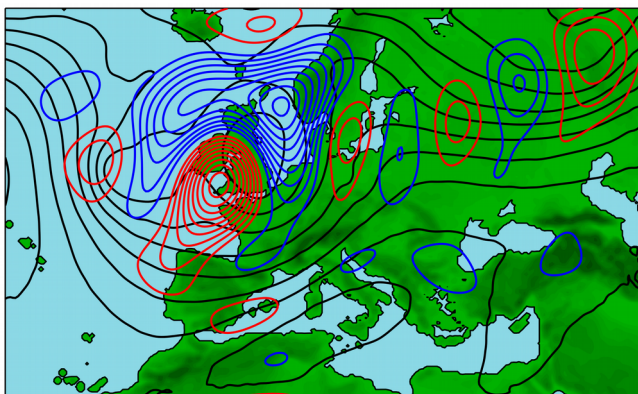
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28 km resolution,
Plant-Craig stochastic conv.

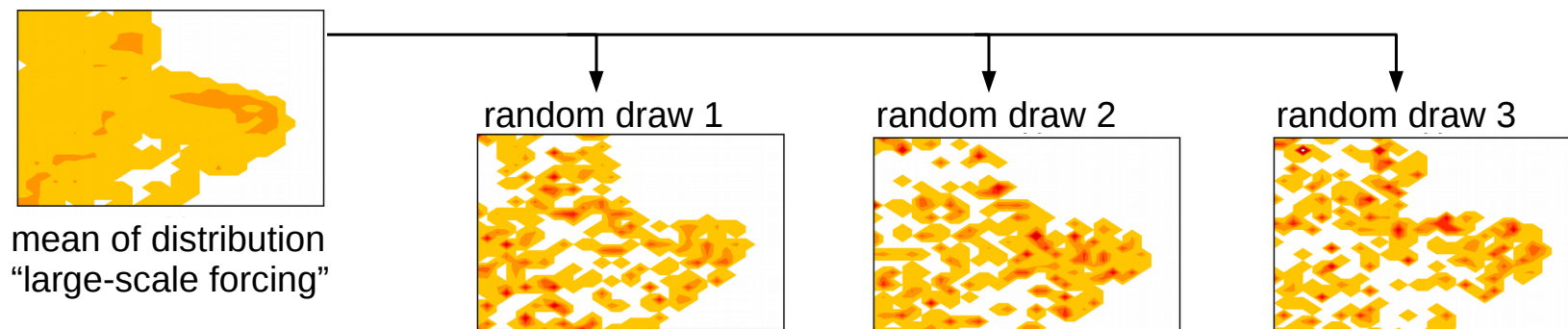


- Conventional convection schemes do not amplify errors near the convective scale sufficiently → **overconfidence**
- The Plant-Craig stochastic convection scheme showed **similar errors** than the convection-permitting reference run

Errors in 500hPa geopot after 60h (color)

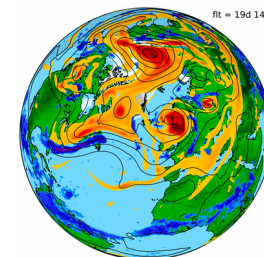
Plant-Craig scheme: basic idea

- Closure assumption determines the **mean of a distribution**
- Clouds are **drawn randomly** from this distribution
- **Ensemble of different realizations** (microstates) consistent with the large-scale forcing can be generated

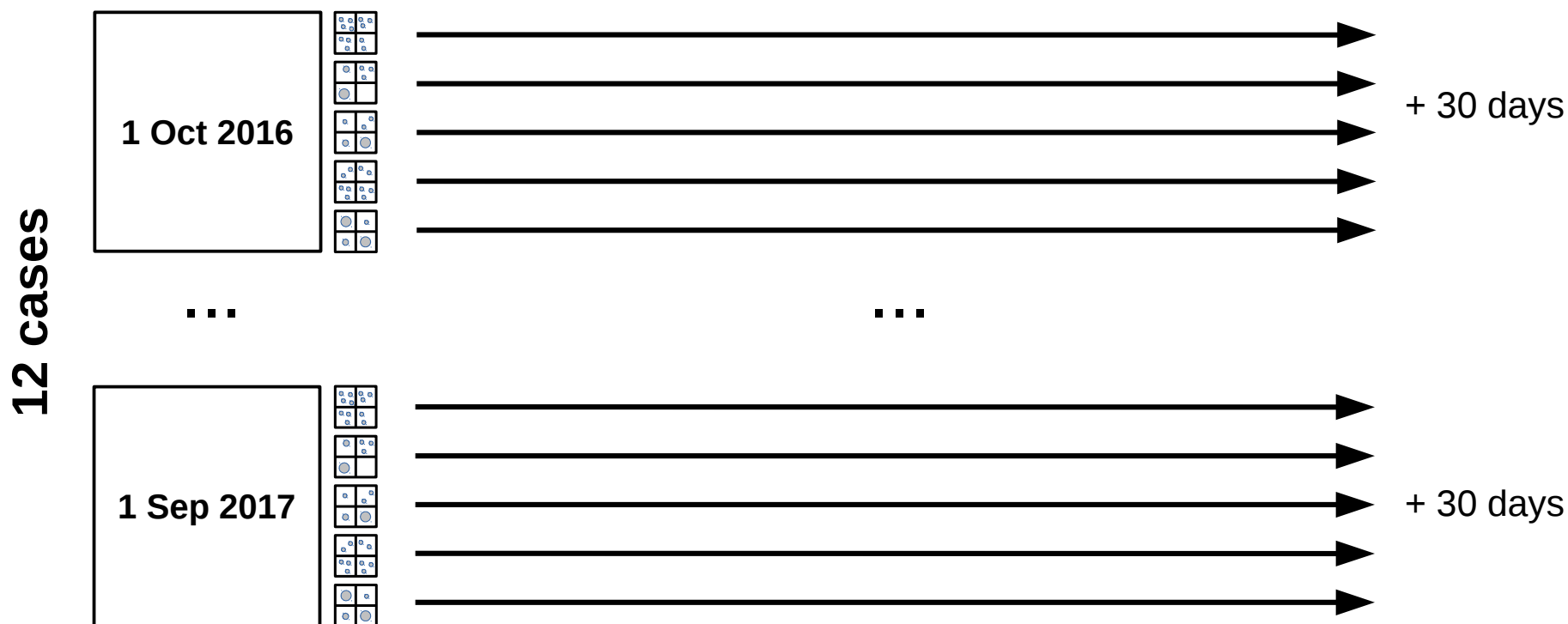


Experimental setup

Experimental - setup



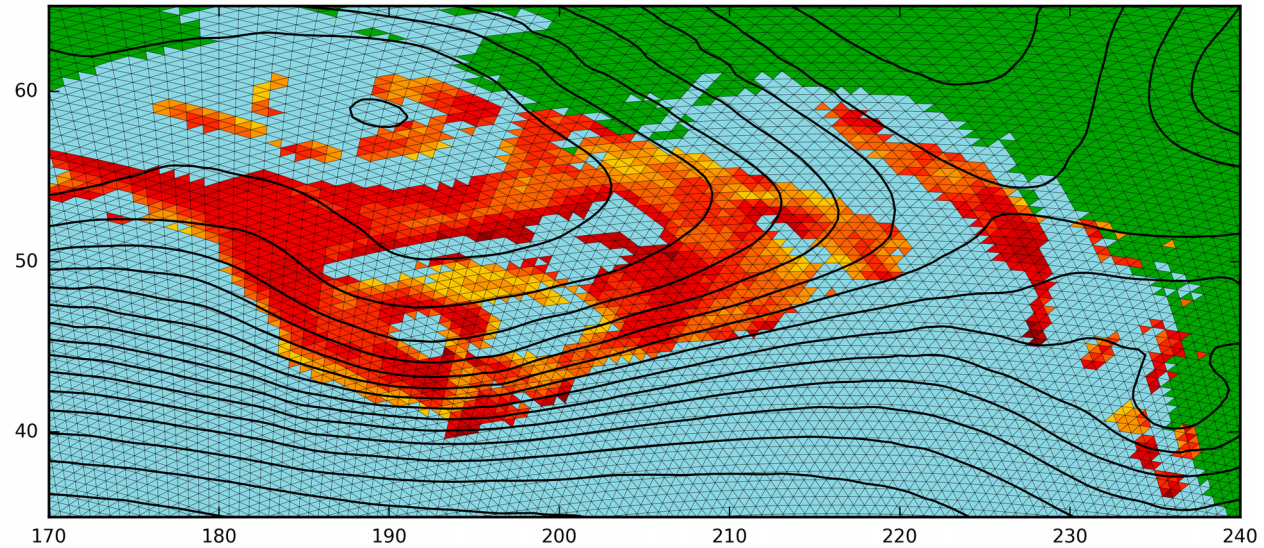
- **Global ICON** simulations (40km resolution)
- **30 days** forecast lead time
- **12 recent cases à 5 members**
- **Plant-Craig convection scheme** to estimate convective-scale uncertainty
- IFS ensemble (50 members) as reference for **current forecasting abilities**



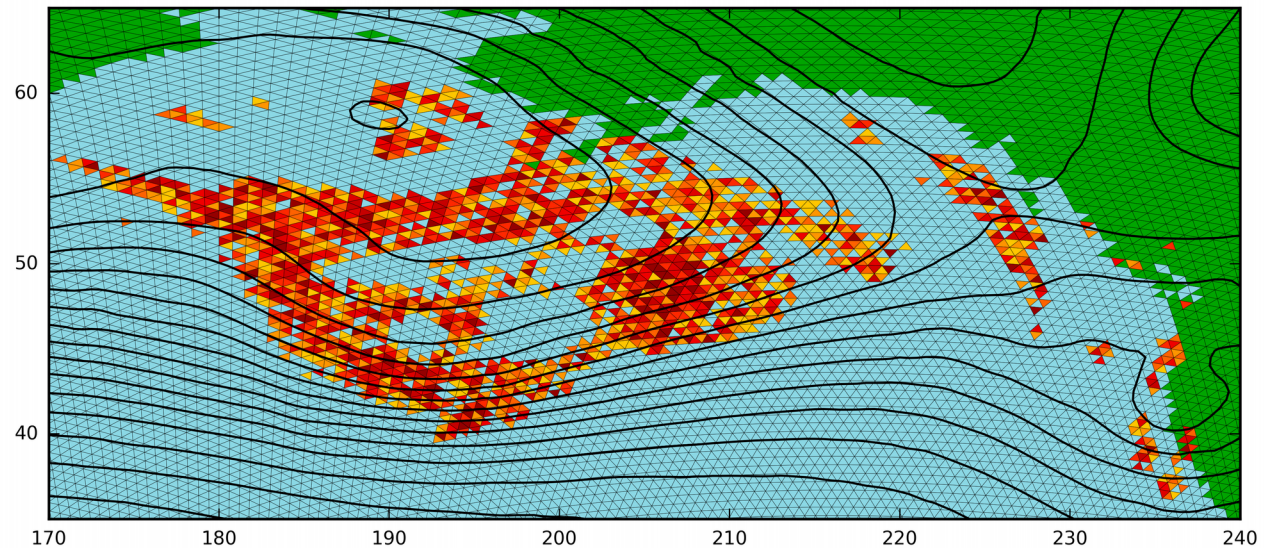
Results

Example: 1 Nov 2016, 01UT, Eastern North Pacific

Closure massflux

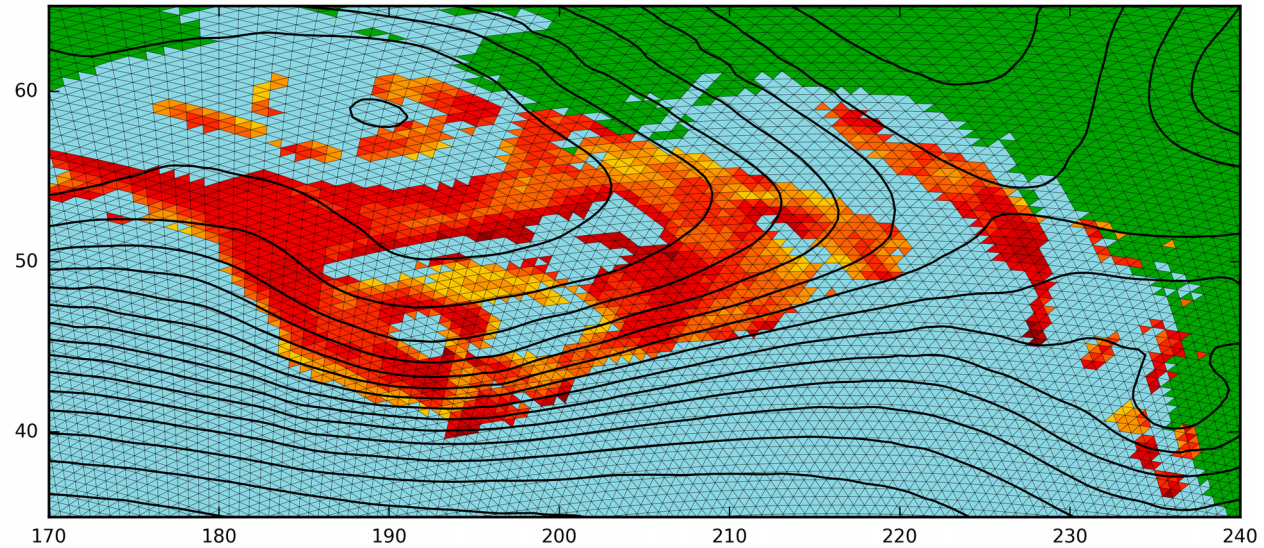


Realized massflux,
Member #1

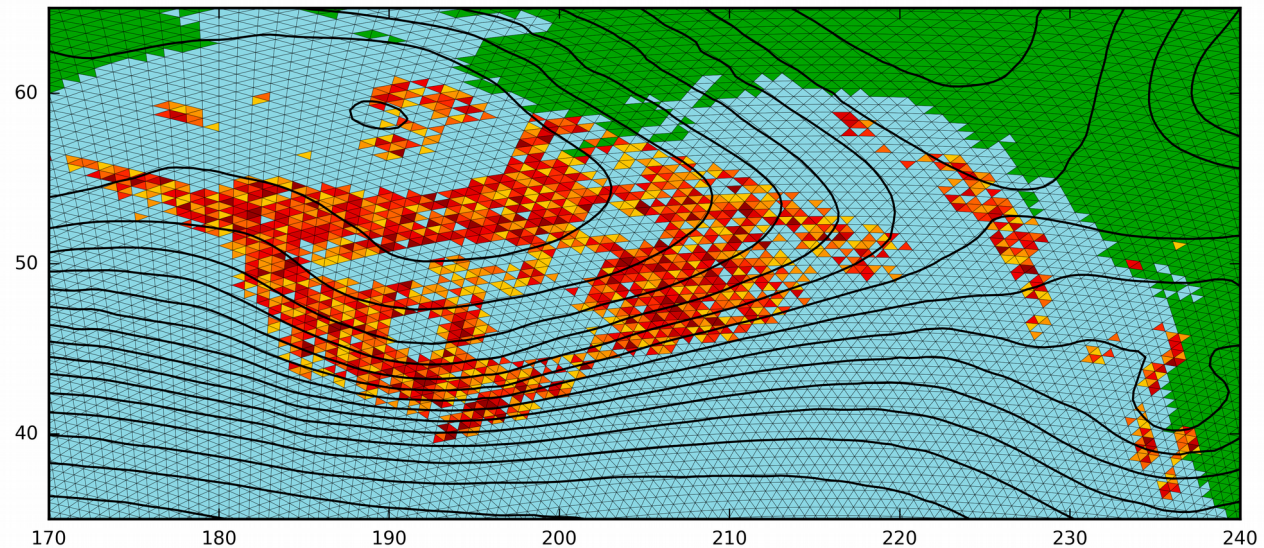


Example: 1 Nov 2016, 01UT, Eastern North Pacific

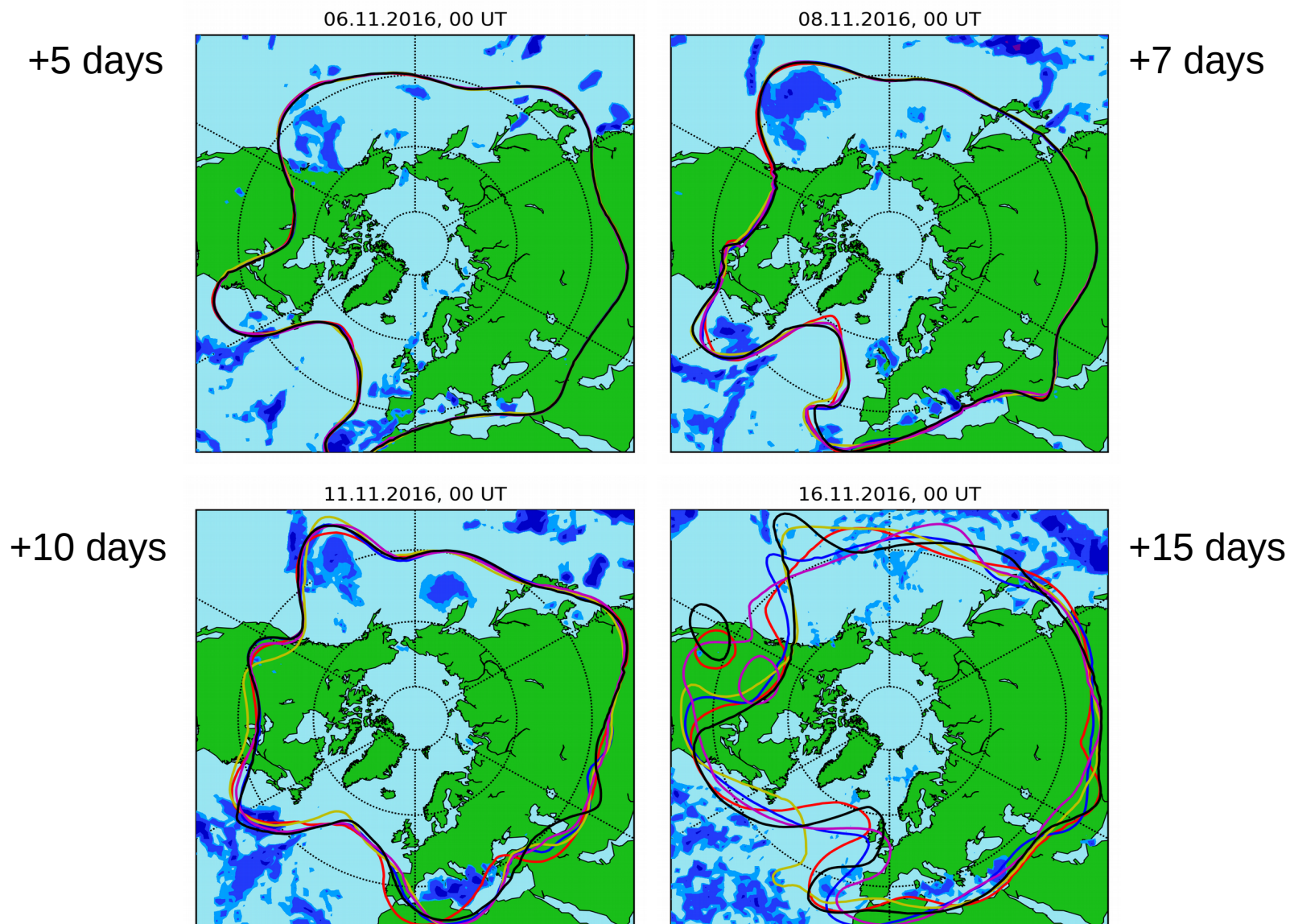
Closure massflux



Realized massflux,
Member #2



Example: 1 Nov 2016-run, 300 hPa geopotential



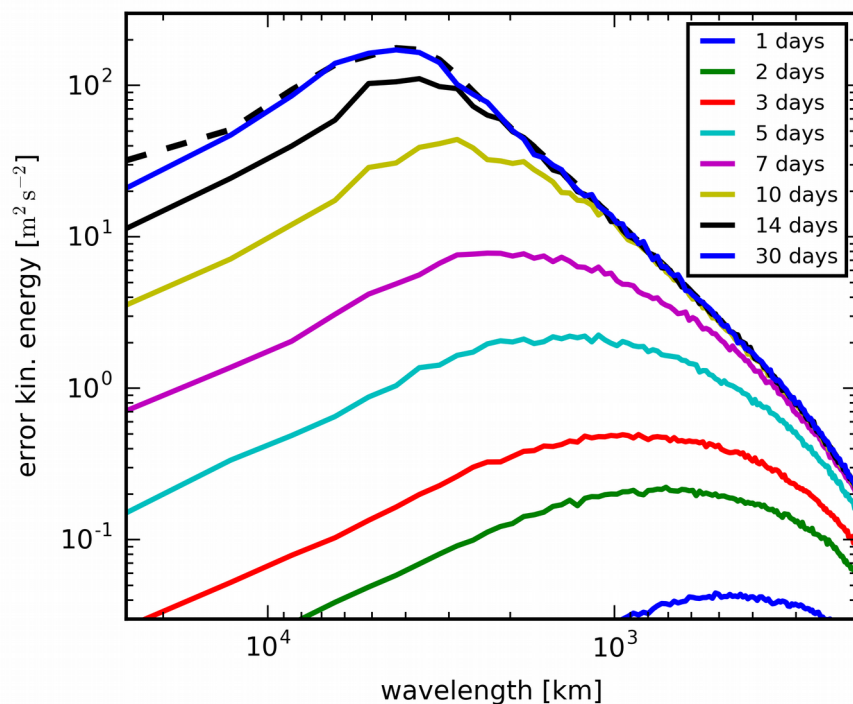
Mid-latitude spectral error kinetic energy (EKE) at 300hPa

- Only mid-latitudes (40° - 60°)
- Average over all 12 cases
- Average over both hemispheres

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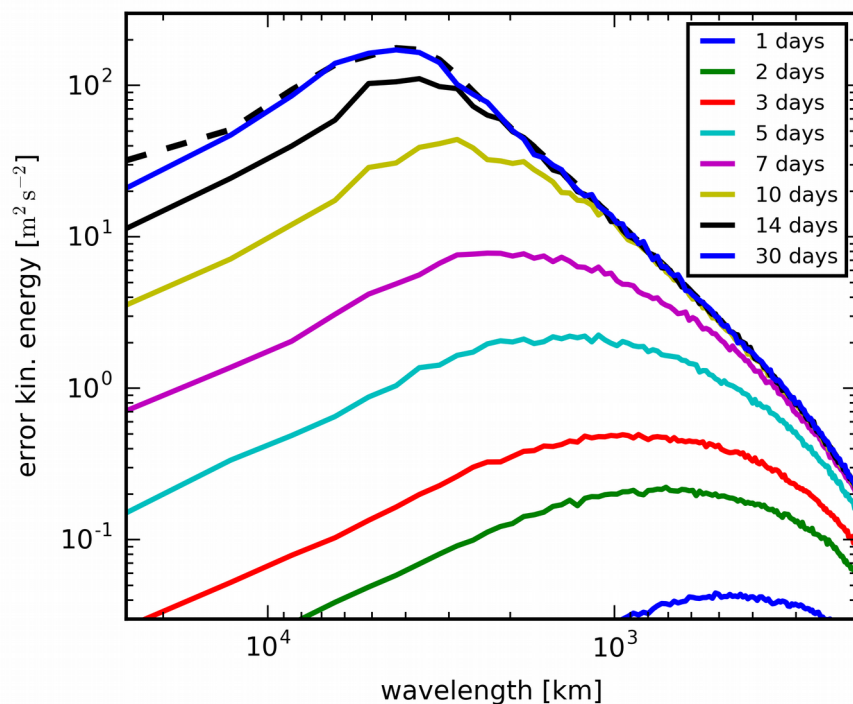
PC-Ensemble



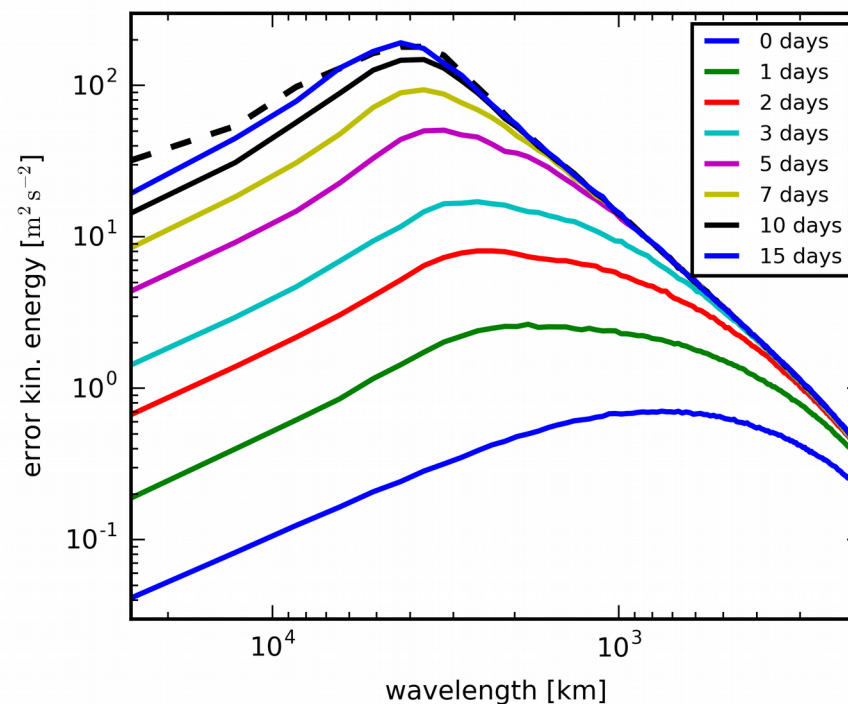
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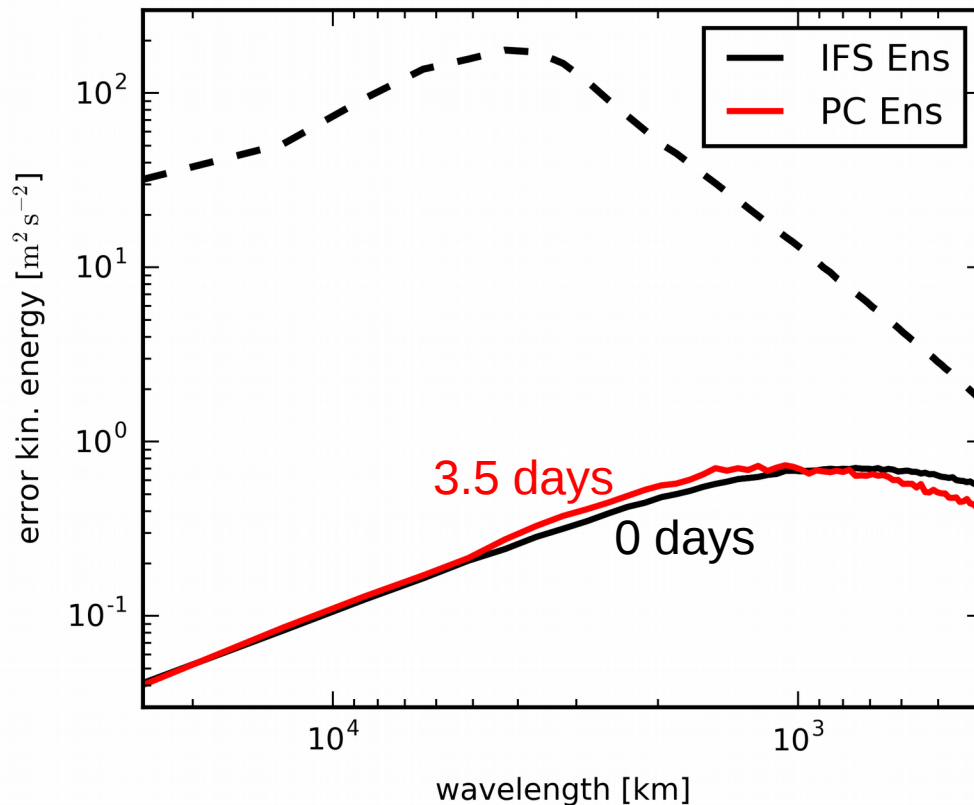
PC-Ensemble



IFS-Ensemble

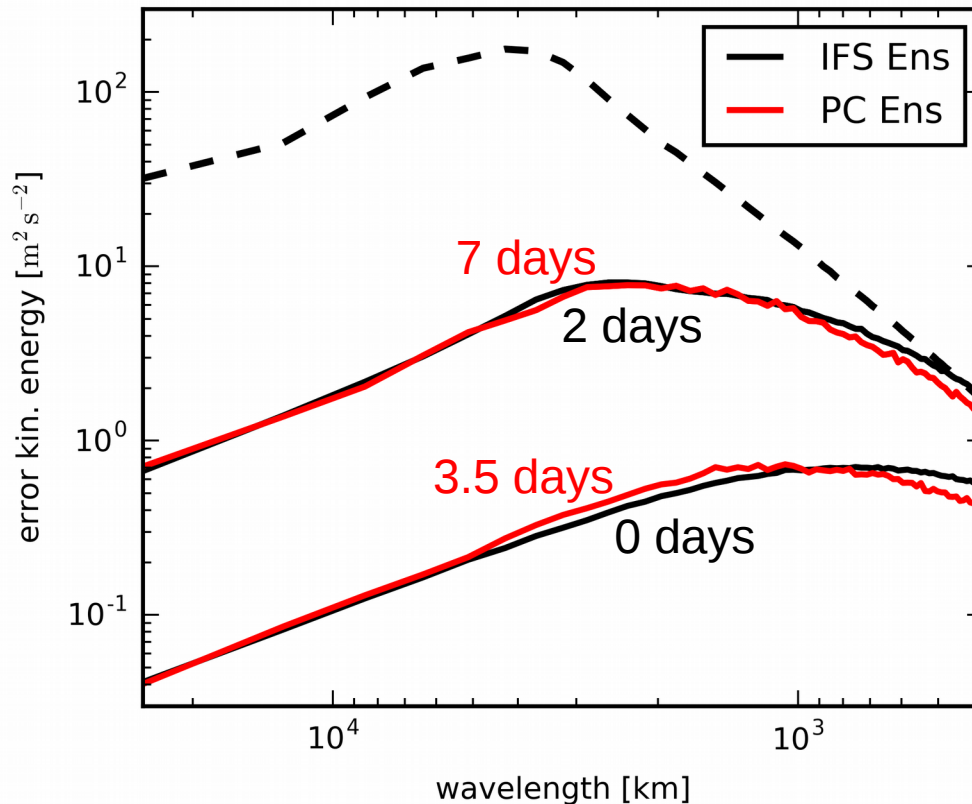


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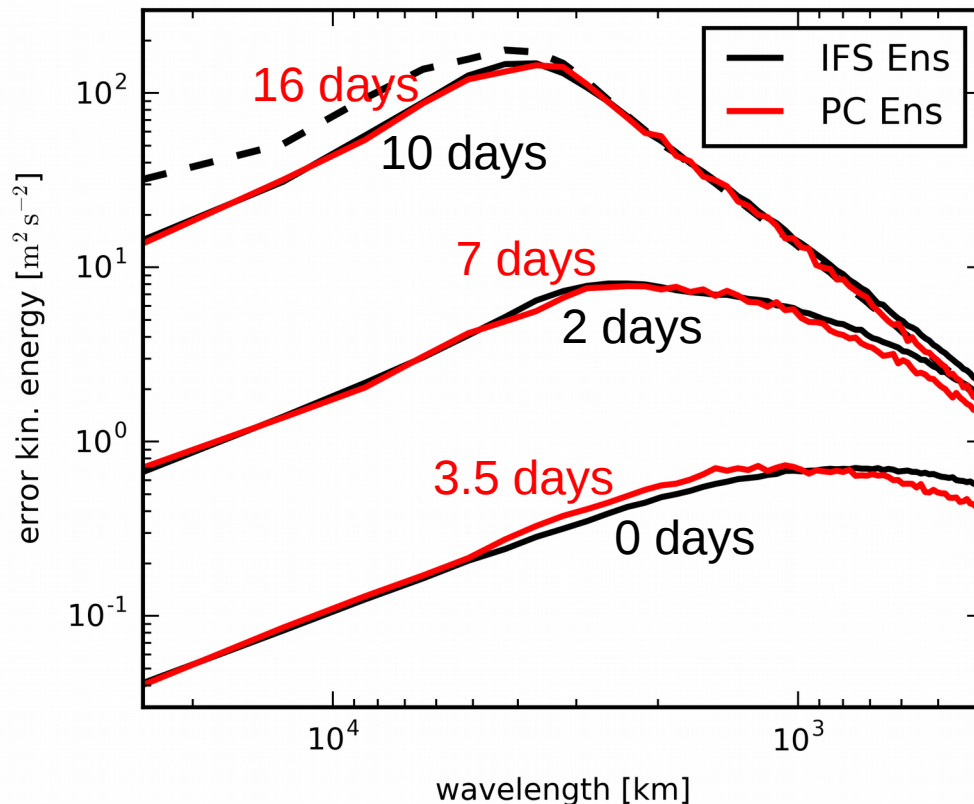
- IFS initial condition uncertainty compares to 3.5 days of upscale error growth

Mid-latitude error kinetic energy (EKE) at 300hPa



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- IFS error grows faster (inflation by singular vectors and SPPT)

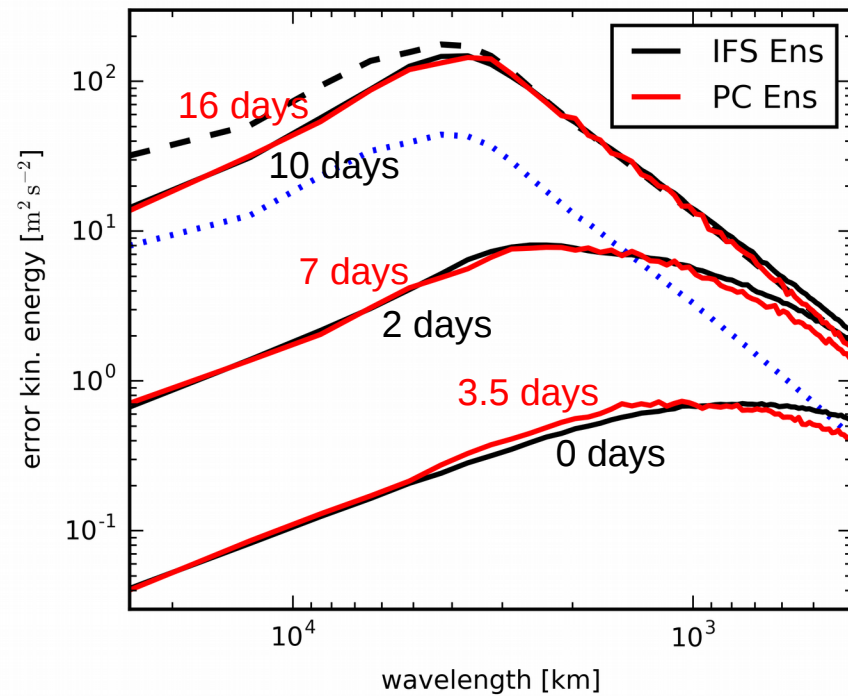
Mid-latitude error kinetic energy (EKE) at 300hPa



- IFS initial condition uncertainty compares to 3.5 days of upscale error growth
- IFS error grows faster (inflation by singular vectors and SPPT)
- Time gap extends to ca. 6 days

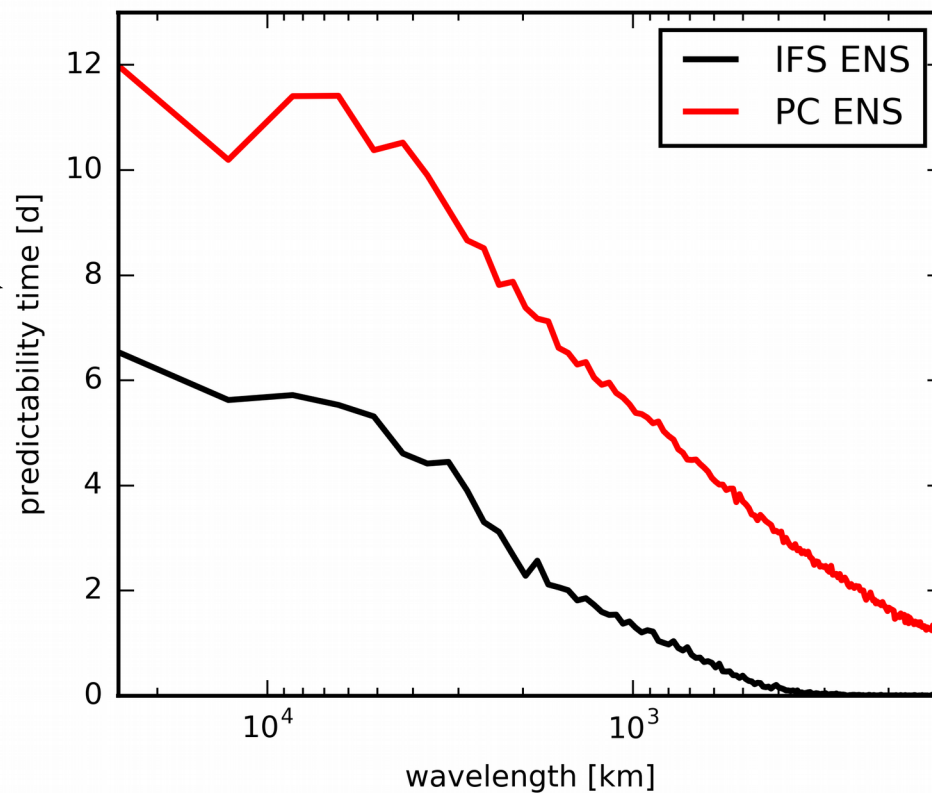
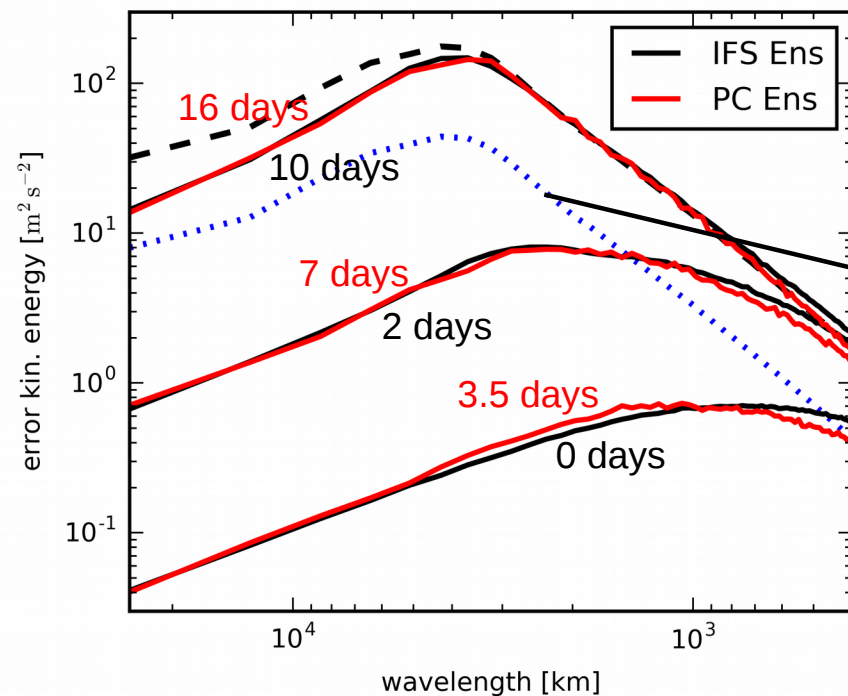
Predictability time from 25% threshold

Define a threshold (25%):



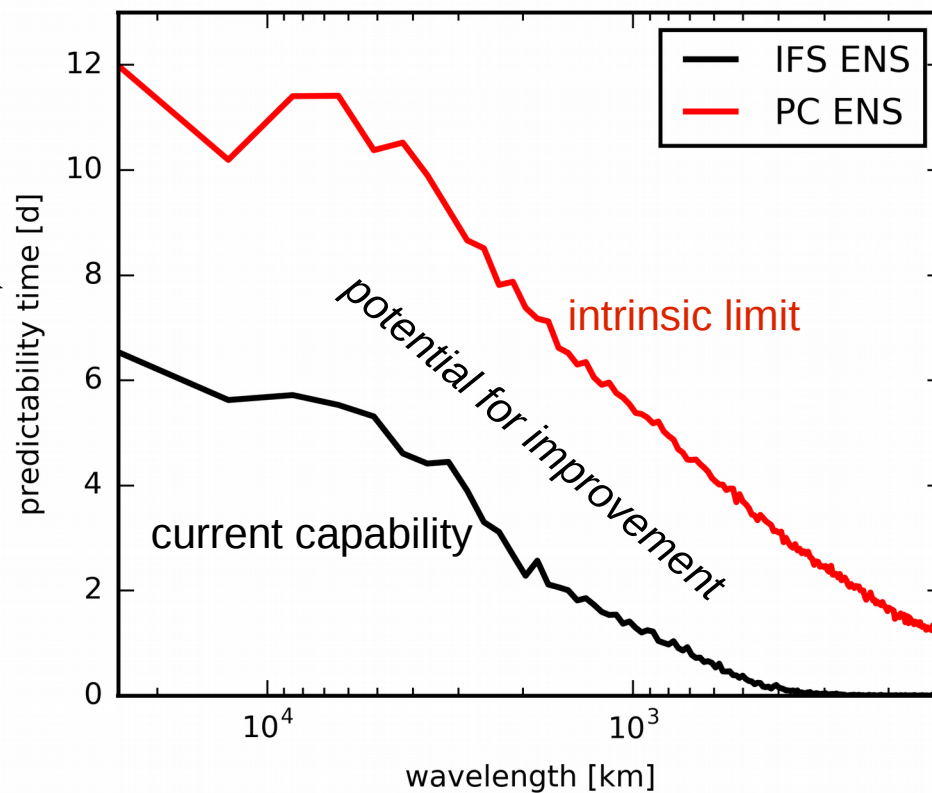
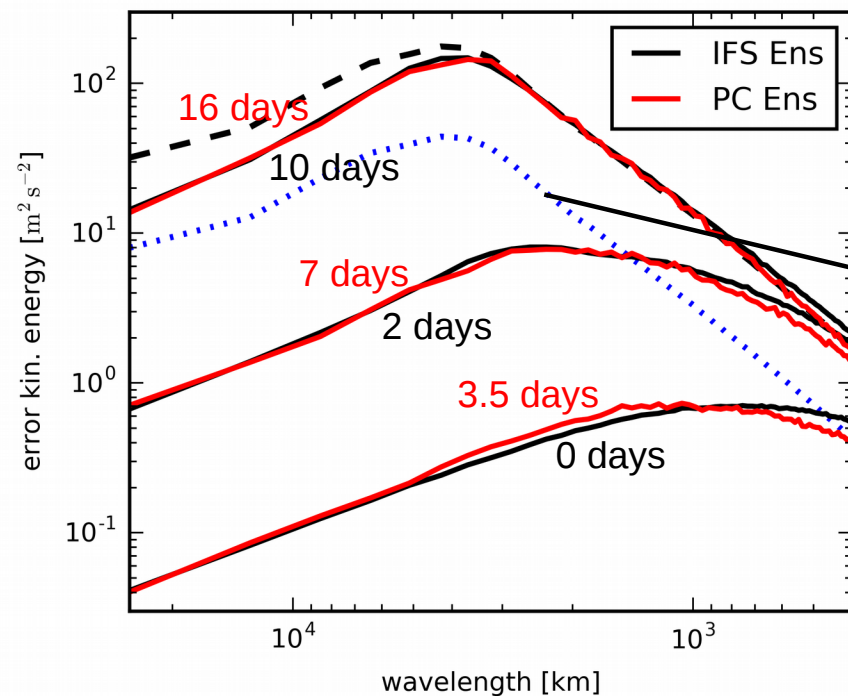
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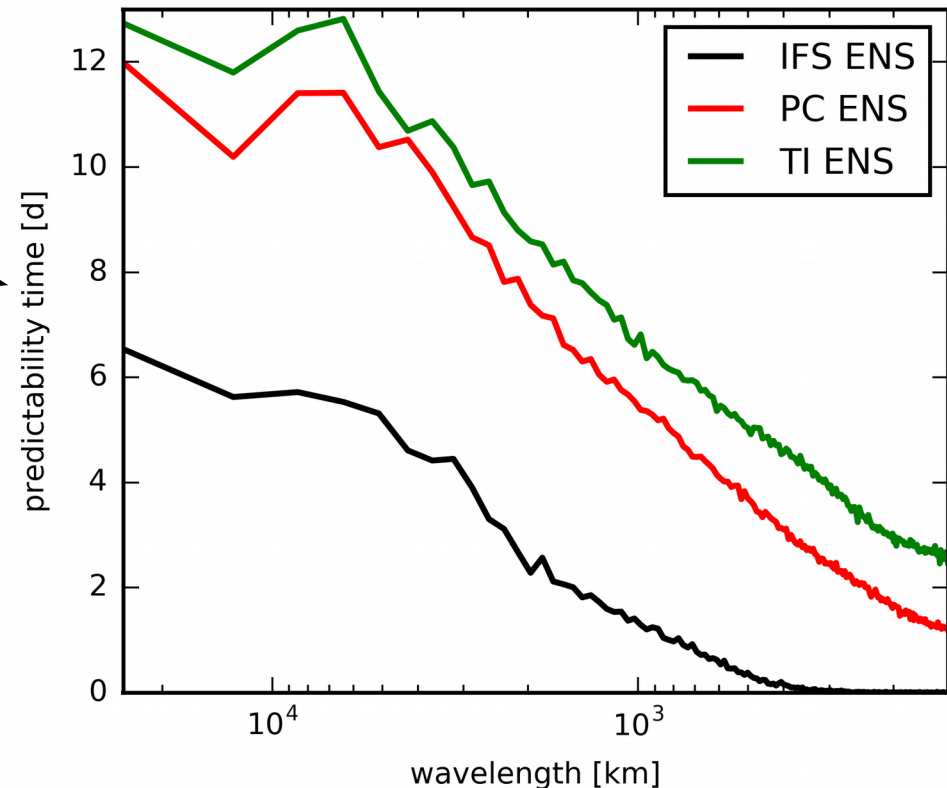
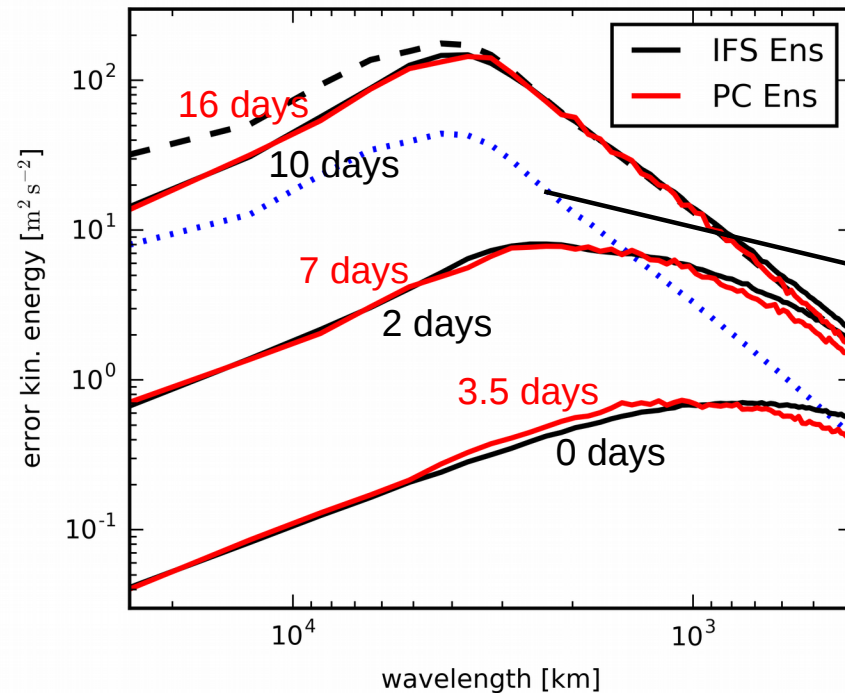
Predictability time from 25% threshold

Define a threshold (25%):



Predictability time from 25% threshold

Define a threshold (25%):



- Tiedtke scheme gives longer intrinsic predictability estimates (overconfidence)
- Difference gets smaller for large modes and long predictability times

Conclusions

- **Upscale propagation time** from convective scale to planetary scale has been estimated to **around 15-20 days**
- The error growth in the PC-ensemble estimates the **intrinsic predictability limit** since predictability of convection cannot be extended beyond its intrinsic limit of $O(10 \text{ hours})$
- Forecasts of current ECMWF forecasting system can be improved by **6 days** for the largest scales:
 - $\approx 3.5 \text{ days}$ through **perfecting the initial conditions**
 - $\approx 2.5 \text{ days}$ through **perfecting the model**
- The **Tiedtke** convection scheme **overestimates** the intrinsic predictability at Mesoscale and synoptic scale but not (much) at planetary scale

Definition of spectral Error Kinetic Energy (similar for v):

$$\begin{aligned} \frac{1}{2} |\tilde{u}_1 - \tilde{u}_2|^2 &\longrightarrow \frac{1}{N^2 - N} \sum_{i \neq j} \frac{1}{2} |\tilde{u}_i - \tilde{u}_j|^2 \\ &= \frac{1}{N-1} \sum_i |\tilde{u}_i - \bar{\tilde{u}}|^2 = \frac{N}{N-1} \left(\frac{1}{N} \sum |\tilde{u}_i|^2 - \left| \frac{1}{N} \sum \tilde{u}_i \right|^2 \right) \end{aligned}$$